Climate Change – What we know and what not

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(Note: due to unclear copyright situations, some of the figures have been removed from the original presentation, but references are given)



Content



• Basics

- What is climate?
- Atmospheric energy balance
- The current state of climate research
 - Observations
 - Causes of climate change
 - Models
 - Projections



Weather and Climate



"Climate is what you expect, weather is what you get."

(Robert A. Heinlein, 1907-1988, amerikanischer Schriftsteller)





"The word climate, however, denotes first and foremost a specific property of the atmosphere, but this property depends on the perpetual interactions of a fully and deeply moving, of currents of contrasting temperatures furrowed sea surface with the heat-radiating dry Earth, which is manifoldly structured, elevated, coloured, naked, or covered with woods and herbs."

A. v. Humboldt, "Kosmos – Entwurf einer physischen Weltbeschreibung ", 1845-1862, translated by H. Schmidt (Picture of A. v. Humboldt)



Climate as a probability distribution



(time series and pdf of some climate variable)



The atmospheric energy budget



(Fig. 7 from Kiehl and Trenberth: Earth's annual global mean energy budget, Bulletin of the American Meteorological Society, 1997)





IPCC: Intergovernmental Panel on Climate Change

Established in 1988 by WMO und UNEP

Organized in 3 working groups:

- WG I assesses the physical scientific basis of the climate system and climate change.
- WG II assesses the vulnerability of socio-economic and natural systems to climate change.
- WG III assesses options for mitigating climate change.

About every 6 or 7 years reports of the working groups are published. Hundereds of scientists are involced in this exercise. The "Fifth Assessement Report" (AR5) was published in 2013/14. The publication series of AR6 reports will start on August, 9, 2021 with WG1.

Downloads: www.ipcc.ch



The current state of climate research



WORKING GROUP I CONTRIBUTION TO THE IPCC AR5 CLIMATE CHANGE 2013: THE PHYSICAL SCIENCE BASIS

(http://www.ipcc.ch)



INTERGOVERNMENTAL PANEL ON CLIMATE CHANGE

IPCC Special Report on Global Warming of 1.5 °C

(published in 2018)

(http://www.ipcc.ch)

ipcc

INTERGOVERNMENTAL PANEL ON CLIMATE CHANES

Global Warming of 1.5°C

An IPCC Special Report on the impacts of global warming of 1.5°C above pre-industrial levels and related global greenhouse gas emission pathways, in the context of strengthening the global response to the threat of climate change, sustainable development, and efforts to eradicate poverty



WORKING GROUP I CONTRIBUTION TO THE **IPCC AR6** (expected to be published August, 9, 2021)





Uncertainties – The language of the IPCC



Agreement	High agreement Limited evidence	High agreement Medium evidence	High agreement Robust evidence			
	Medium agreement Limited evidence	Medium agreement Medium evidence	Medium agreement Robust evidence	Confidence Scale		
	Low agreement Limited evidence	Low agreement Medium evidence	Low agreement Robust evidence			

Evidence (type, amount, quality, consistency)

The following terms have been used to indicate the assessed likelihood, and typeset in italics:

Term*	Likelihood of the outcome
Virtually certain	99–100% probability
Very likely	90–100% probability
Likely	66–100% probability
About as likely as not	33–66% probability
Unlikely	0–33% probability
Very unlikely	0–10% probability
Exceptionally unlikely	0–1% probability

* Additional terms (*extremely likely*: 95–100% probability, *more likely than not*: >50–100% probability, and *extremely unlikely*: 0–5% probability) may also be used when appropriate.





Observed climate change



(NOAA National Centers for Environmental information, Climate at a Glance: Global Time Series, published July 2021, retrieved on July 18, 2021 from <u>https://www.ncdc.noaa.gov/cag/</u>)



(NOAA National Centers for Environmental information, Climate at a Glance: Global Time Series, published July 2021, retrieved on July 18, 2021 from <u>https://www.ncdc.noaa.gov/cag/</u>)





(Fig. 7 from Kiehl and Trenberth: Earth's annual global mean energy budget, Bulletin of the American Meteorological Society, 1997)

Increased downward infrared radiation at the surface has to be compensated by surface warming which leads to increased longwave emission of the surface and increased latent and sensible heat transfer. However, evaporation is limited over dry surfaces.





Observed regional temperature increase



(b) Observed change in average surface temperature 1901–2012



(IPCC AR5, SPM, 2013)

Observations: Arctic sea ice, global sea level





(IPCC AR5, SPM, 2013)





" The observed linear relationship implies a sustained loss of 3 \pm 0.3 square meters of September sea-ice area per metric ton of CO₂ emission. "

Notz and Stroeve, Science, 2016

(September Arctic sea ice figures for 1984 and 2016 from nsidc.org)





Observed change in precipitation over land

1901-2010

1951-2010



(IPCC AR5, SPM, 2013)





(Fig. 1a from O'Gorman: Precipitation Extremes Under Climate Change, Current Climate Change Reports, 2015) Change of annual maximum of daily precipitation per global temperature increase of 1K





Causes for climate change



Natural forcings

- Volcanic eruptions
- Variations of solar activity
- Variations of Earth's orbital parameters (Milankovic cycles)

Anthropogenic forcings

- Fossil fuel burning (coal, oil, gas) (e.g., CO₂, SO₂, soot)
- Methane emissions (e.g. from fossil fuel production and transport, livestock farming, landfills)
- Land use changes (e.g. deforestation, urbanisation)

Attention, this list is by no means complete ... and there is internal variability of the climate system



The eruption of Mount Pinatubo



In the year after the eruption of Mt. Pinatubo, global surface temperature decreased by about 0.4°C.





400 Years of Sunspot Observations



(source: Wikipedia)

CO₂ in the atmosphere





(IPCC AR5, SPM, 2013)



(www.co2.earth)



Greenhouse gases in ice cores





(IPCC, 2007)

Svante Arrhenius (1859-1927)





THE LONDON, EDINBURGH, AND DUBLIN PHILOSOPHICAL MAGAZINE AND JOURNAL OF SCIENCE. [FIFTH SERIES.] APRIL 1896.

XXXI. On the Influence of Carbonic Acid in the Air upon the Temperature of the Ground. By Prof. SVANTE ARRHENIUS *.

One may now ask, How much must the carbonic acid vary according to our figures, in order that the temperature should

See Tyndall Lecture by Raymond Pierrehumbert, AGU, 2012: https://www.youtube.com/watch?v=RICBu_P8JWl

Or read https://history.aip.org/climate/Radmath.htm

4°-5° C.). The demands of the geologists, that at the genial

Radiative forcings of the climate system



	Emitted Compound	Resulting Atmospheric Drivers	Radia	ative Forcin	ng by Emiss	ions and D	Drivers	Level of Confidence	
Anthropogenic	CO ²	CO ₂				• <u>+</u>	1.68 [1.33 to 2.03]	VН	
	CH4	CO ₂ H ₂ O ^{str} O ₃ CH ₄				1	0.97 [0.74 to 1.20]	н	
	Halo- carbons	O ₃ CFCs HCFCs				1	0.18 [0.01 to 0.35]	н	
	N ₂ O	N ₂ O			l.		0.17 [0.13 to 0.21]	VH	
	S	CO ₂ CH ₄ O ₃		I		1	0.23 [0.16 to 0.30]	м	
	NMVOC	CO ₂ CH ₄ O ₃		H		1	0.10 [0.05 to 0.15]	м	
	Gases ar ON	Nitrate CH ₄ O ₃		•			0.15 [-0.34 to 0.03]	м	
	Aerosols and precursors (Mineral dust,	Mineral Dust Sulphate Nitrate Organic Carbon Black Carbon				-	0.27 [-0.77 to 0.23]	н	
	SO ₂ , NH ₃ , Organic Carbon and Black Carbon	Cloud Adjustments due to Aerosols				-0	0.55 [-1.33 to -0.06]	L	
		Albedo Change due to Land Use		H		-0	0.15 [–0.25 to –0.05]	м	
Natural		Changes in Solar Irradiance		•	1	1	0.05 [0.00 to 0.10]	м	
Total Anthropogenic RF relative to 1750			20	11		1	2.29 [1.13 to 3.33]	н	
			19	80			1.25 [0.64 to 1.86]	н	
			19	50		1	0.57 [0.29 to 0.85]	М	
			-1	0	1	2	3		
	Radiative Forcing relative to 1750 (W m ⁻²)								





Climate models



Radiative-Convective Equilibrium



(Fig. 7 from Kiehl and Trenberth: Earth's annual global mean energy budget, Bulletin of the American Meteorological Society, 1997)



Manabe & Wetherald (JAS, 1967) estimated a warming of 2,7 K for a doubling of the atmospheric CO₂ concentration. We (Kluft et al., J. Climate, 2019) repeated their calculations, arriving at 2,8 K.





$$\frac{\partial \mathbf{v}}{\partial t} = -(f+\xi)\mathbf{k} \times \mathbf{v} - \nabla K - \dot{\eta}\frac{\partial \mathbf{v}}{\partial \eta} - \frac{R_d T}{p}\nabla p - \nabla \phi + \mathbf{F}_{\mathbf{v}}$$

$$\frac{\partial T}{\partial t} = -\mathbf{v} \cdot \nabla T - \dot{\eta}\frac{\partial T}{\partial \eta} + \frac{R_d T}{C_p}\frac{dp}{p}\frac{dp}{dt} + F_T$$

$$0 = \frac{\partial \phi}{\partial \eta} + \frac{R_d T}{p}\frac{\partial p}{\partial \eta}$$

$$\frac{\partial}{\partial t}\left(\frac{\partial p}{\partial \eta}\right) = -\nabla \cdot \left(\mathbf{v}\frac{\partial p}{\partial \eta}\right) - \frac{\partial}{\partial \eta}\left(\dot{\eta}\frac{\partial p}{\partial \eta}\right) + F_{\rho}$$

$$\frac{\partial}{\partial t}\left(c_i\frac{\partial p}{\partial \eta}\right) = -\nabla \cdot \left(\mathbf{v}c_i\frac{\partial p}{\partial \eta}\right) - \frac{\partial}{\partial \eta}\left(\dot{\eta}c_i\frac{\partial p}{\partial \eta}\right) + F_{c_i}$$

$$(i = 1, 2, 3, ...)$$

 ξ : vorticity; K: kinetic energy; η : vertical coordinate; c_i : mixing ratio of tracer i F: external forcing. Other symbols have their conventional meanings.

can't be solved analytically ...





... but numerically.



However, sub-grid scale processes need to be parameterized.





(Picture of trees)







(IPCC AR5, WG1, Fig. 9.8, 2013)







Observations

Models using only natural forcings Models using both natural and anthropogenic forcings

(IPCC AR5, SPM, 2013)



Climate sensitivity



Climate sensitivity: Temperature increase ΔT resulting from a radiative forcing ΔF at the tropopause.

Usually defined as Temperature increase for a doubling of the atmospheric CO₂ concentration (for which a radiative forcing of $\Delta F \approx 3.7 \text{ W/m}^2$ is assumed).

According to IPCC AR5, equilibrium climate sensitivity (ECS) is likely between 1.5 and 4.5 °C.

A recent study (Sherwood et al., Rev. of Geophysics, 2020) has narrowed this to between 2.6 and 3.9 °C.

Several recent global climate models have sensitivities above these ranges.

I'm curious to see the upcoming IPCC assessment on this topic.

(IPCC AR5, WG1, TFE.6, Fig.1, 2013)



One important reason for differences in climate sensitivity among models?











Projections of future climates





(Fig. 3 from van Vuuren et al., The representative concentration pathways: an overview Climatic Change, 2011)

Green: RCP2.6 Red: RCP4.5 Black: RCP6.0 Blue: RCP8.5

RCP8.5 close to SRES A2, RCP4.5 close to SRES B1

Projections: global temperature, Arctic sea ice





Cumulative CO₂-emissions and global temperature





(IPCC AR5, SPM, 2013)





Remaining budget (from 2015):

650±130 Gt(CO₂) ΔT<1.5°C: 1300±130 Gt(CO₂) **ΔT<2°C**:

Current emission rate:

~40 Gt(CO₂)/yr

(Lawrence et al., Nature Communications, 2018)

How much can we still emit?



Global warming relative to 1850-1900 (°C)



"In model pathways with no or limited overshoot of 1.5C, global net anthropogenic CO2 emissions decline by about 45% from 2010 levels by 2030 (40–60% interquartile range), reaching net zero around 2050 (2045–2055 interquartile range). For limiting global warming to below 2C CO2 emissions are projected to decline by about 25% by 2030 in most pathways (10–30% interquartile range) and reach net zero around 2070 (2065–2080 interquartile range). Non-CO2 emissions in pathways that limit global warming to 1.5Cshow deep reductions that are similar to those in pathways limiting warming to 2C."

...

"Most 1.5 °C and 2 °C pathways are heavily reliant on CDR (Carbon Dioxide Removal) at a speculatively large scale before midcentury."

> (IPCC, 2018: Global Warming of 1.5°C; Fig. SPM.1; SPM, C.1; Section 2.6.3)



2060

2100

1980

2020

2060

2020

1980

Maximum temperature rise is determined by cumulative net CO₂ emissions and net non-CO₂ radiative forcing due to methane, nitrous oxide, aerosols and other anthropogenic forcing agents.

1980

2020

2060

2100

2100







(Fig. 7 from Kiehl and Trenberth: Earth's annual global mean energy budget, Bulletin of the American Meteorological Society, 1997)

Latent heating from condensation balances tropospheric radiative cooling to space. This cooling (by CO₂ und H₂O) increases under current GHG induced warming of 1°C by about 2%. Likewise react latent heating and precipitation. So global precipitation increases less than humidity (about 7%/1°C following Clausius-Clapeyron) under warming. Precipitation is constrained by energy, not humidity.

Precipitation changes in central Europe, summer









Winter mean zonal wind in 850 hPa (gray isolines, 5 m/s) und its change in the 21. century under strong CO_2 increase

(Fig. 4 from Shepherd et al., Atmospheric circulation as a source of uncertainty in climate change projections, Nature Geoscience, 2014)





50-year trends (2010-2060) simulated with the CCSM3 model

(Fig. 10a from Deser et al., Projecting North American Climate over the Next 50 Years: Uncertainty due to Internal Variability J. Climate, 2013)



Projections of different models



(Fig. 1 from Stevens & Bony, What Are Climate Models Missing, Science, 2013)



(One) future of modelling



Tropical Atlantic in ICON, August, 23, 2016

ICON HErZ - NARVAL-II - HD(CP)² Simulations: 20160823 +10.0h



Simulation by Daniel Klocke (DWD) and vizualization by Matthias Brueck (MPI-M)





- Heat waves:
- Extreme precipitation:
- Coastal floodings:
- Droughts:
- Tropical storms:
- Extratropical storms:

- very likely, globally
- very likely in mid-latitude and tropical land areas
- very likely
- likely in sub-tropical regions
- uncertain
 - uncertain

(IPCC AR5, Tab. SPM.1, 2013)





- "It is extremely likely that human influence has been the dominant cause of the observed warming since the mid-20th century." (IPCC AR5)
- A doubling of atmospheric CO₂ likely causes global warming between 2.6 and 3.9 °C (Sherwood et al., 2020; between 1.5 and 4.5 °C., IPCC AR5)
- The temperature increase until the end of this century depends largely on the cumulative GHG emissions. In order to reach the 2 (or 1,5) °C-limit future cumulative CO₂ emissions should not be larger than 1200 (550) Gt CO₂ emittiert werden. In the past, humans have emitted about 2000 Gt, currently about 40 Gt per year.
- Precipitation has likely already decreased in dry regions and increased in wet regions. This trend will very likely continue. (IPCC, AR5) Global mean precipitation is expected to increase by about 2% per 1°C global warming.
- Currently, sea level rises by about 3 mm per year. A melting of Greenland would increase sea level by about 7m of Antarctica by more than 40m.
- Several weather extremes will likely increase.
- The global thermodynamics of climate change are well understood, how warming affects the weather (in particular circulation) less good.
- I expect there will be surprises. Understanding the system better is the best preparation.

Thank you very much !