The long (!) nucleation phase of laboratory earthquakes

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Thanks to:

- @ ENS: Samson Marty, Blandine Gardonio, Harsha Bhat, Raul Madariaga
- @ IGN, Paris VII: Kristel Chanard
- @ IRAP Toulouse: Soumaya Latour
- @ EPFL: François Passelègue
- @ Durham U.: Stefan Nielsen







Established by the European Commission

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Integrating 30 years of experimental rock fracture mechanics: Ohnaka's view

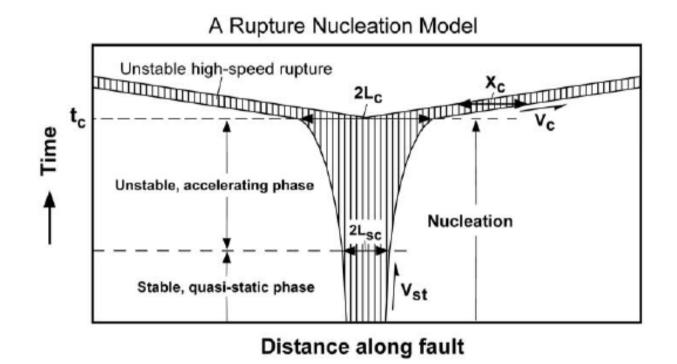


Figure 15. A physical model of rupture nucleation. Hatched portion indicates the zone in which the breakdown (or slip-weakening) proceeds with time. Ohnaka 2003

Integrating 30 years of experimental rock fracture mechanics: Ohnaka's view

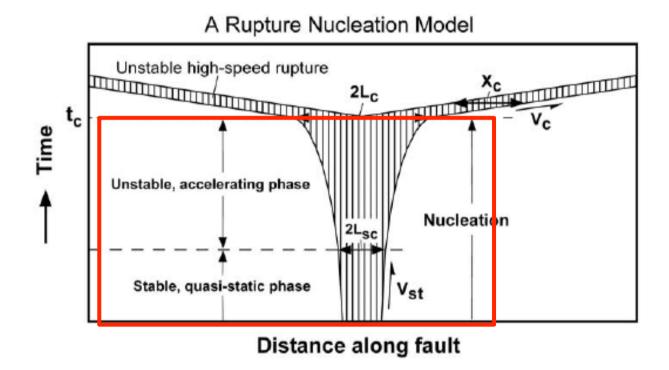


Figure 15. A physical model of rupture nucleation. Hatched portion indicates the zone in which the breakdown (or slip-weakening) proceeds with time. Ohnaka 2003

Sept. 3, 1947

Mr. Herman Saylor 718 1/4 W. 1st Street Los Angeles, California

Dear Sir:

This Laboratory does not predict earthquakes.

Specific predictions iving time and place come from amateurs, publicity-seekers, believers in the occult, or just plain fools.

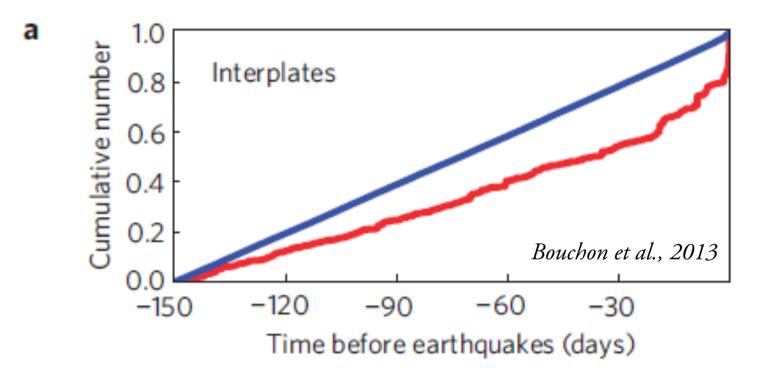
Los Angeles remains exposed to the risk of a great earthquake, which may take place at any time.

Yours truly,

B. Gutenberg Director, Seismological Laboratory

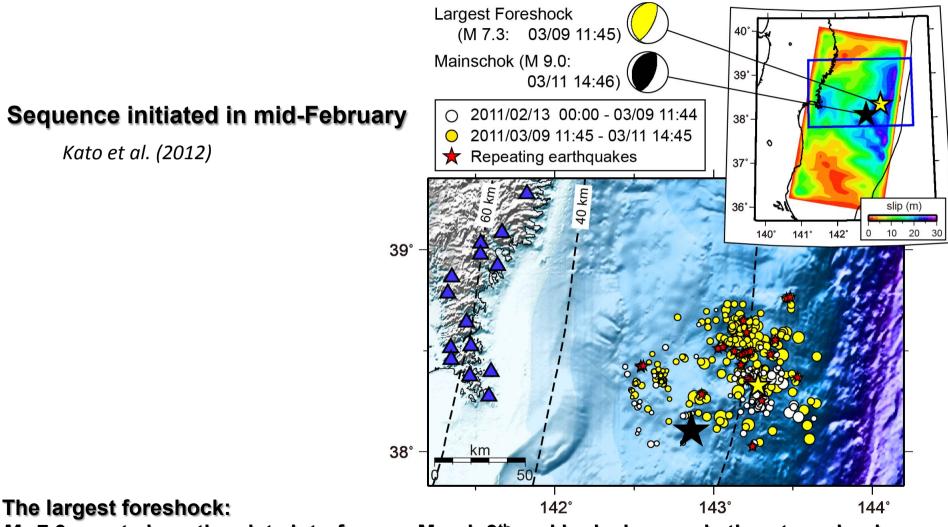
BG:ml

Many earthquakes are preceded by foreshocks



Foreshocks are due to the failure of small asperities prior the main rupture

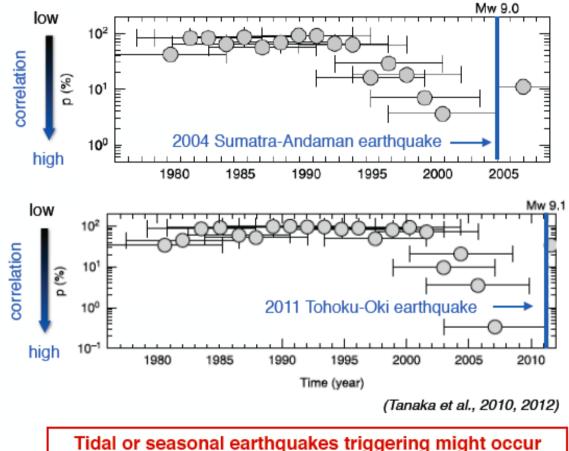
Foreshock sequences before Tohoku-Oki



M_w 7.3 event along the plate interface on March 9th and had a low-angle thrust mechanism.

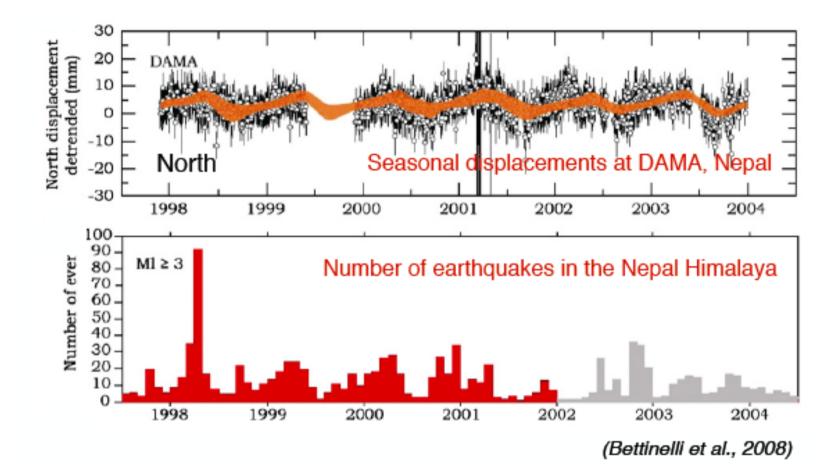
Tidal triggering during large EQ nucleation

Observation: increase of correlation between seismicity and oceanic tides prior to large earthquakes



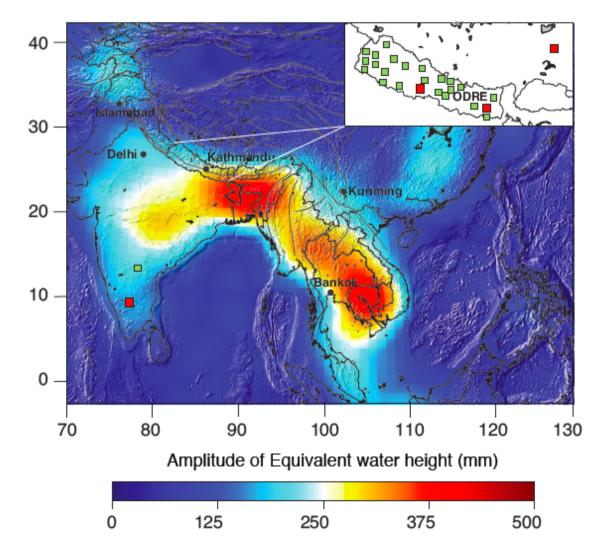
favorably during the long nucleation phase

Seasonal seismicity in the Himalayas



Seasonal seismicity in the Himalayas

Seasonal load in South Asia derived from GRACE (2002-2012) & GPS stations



Outline

- Rock fracture in the lab.
- Precursory processes to Stick-slip
- Foreshock dynamics
- "Tidal" modulation

Quasi-static propagation

View from the lab - quasi static rupture in granite:

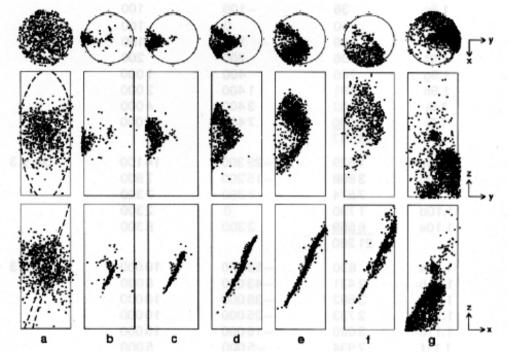
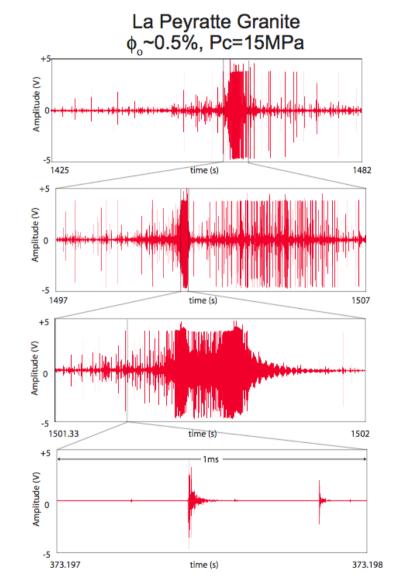


Figure 7. Plots of AE locations for sample G1. Bottom plot of each set is view looking along-strike of eventual fracture plane. Middle plot is face-on view of fault in which sample has been rotated counterclockwise 90°. Top plot is view looking down on sample. Projections of surface trace of eventual fault plane are shown in (a) for reference. Each point represents one AE event. Associated statistics are given in Table 1. Stress intervals for each plot are indicated in Figures 3a and 5a. Fault nucleation occurs in plot (b).

Lockner et al 1991

Failure of granite = the B(rac)enchmark

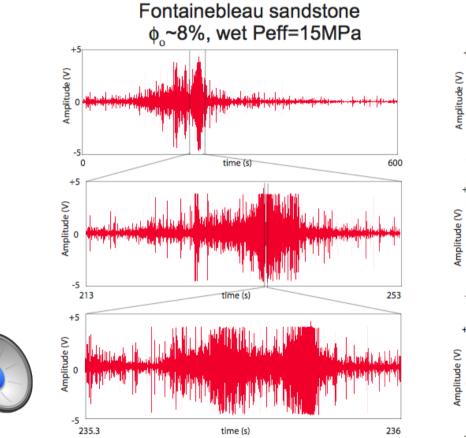


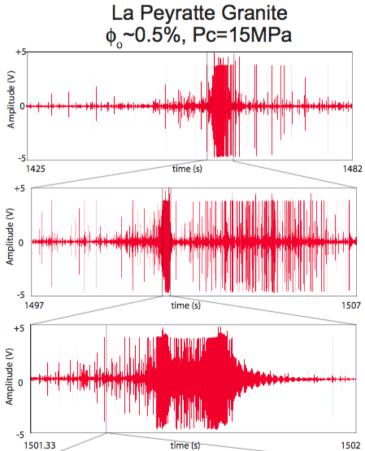
Continuous ultrasonic records of failure (10MHz)

Effect of lithology

Continuous ultrasonic records of failure (10MHz)

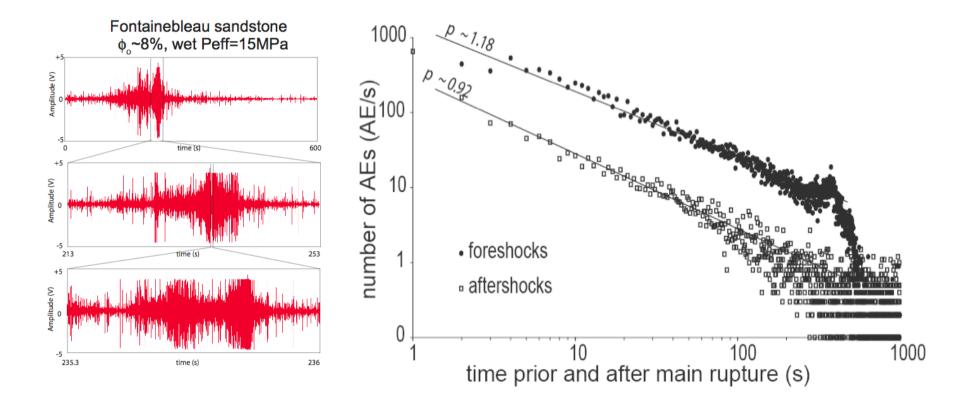
Sandstone vs. granite





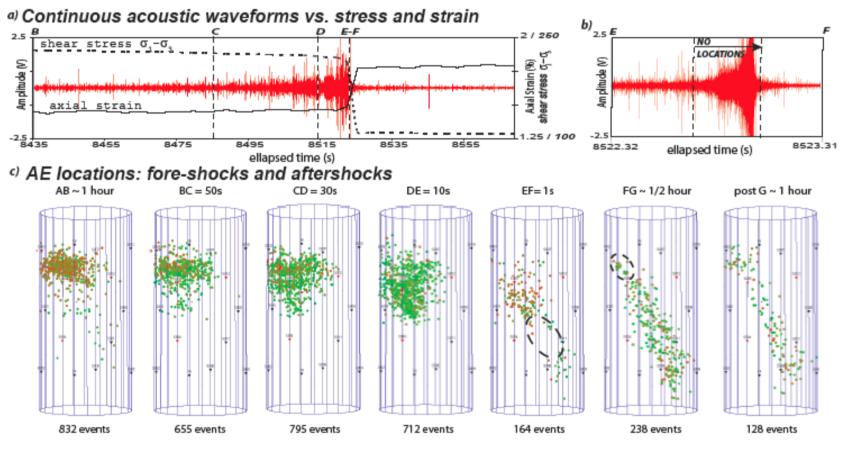
Omori's law

In the lab., foreshocks and aftershocks follow Omori's law, with exponents close to 1



Quasi-dynamic ruptures

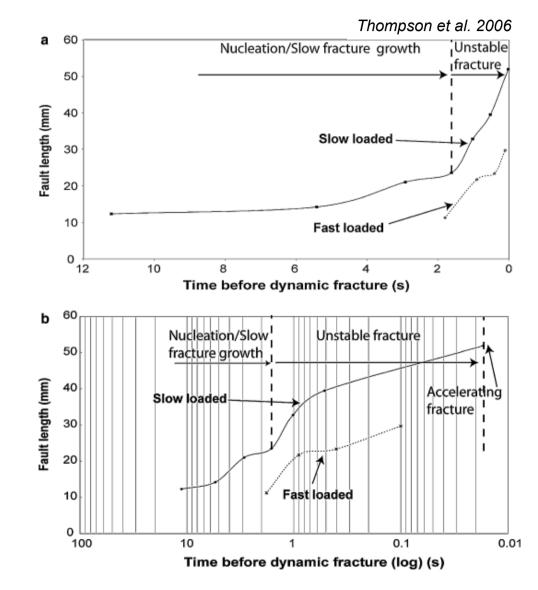
Nucleation zone (~1cm³) Quasi static to dynamic @ >mm/s) Terminal rupture speed (that can be resolved!) <100m/s



Schubnel et al., 2007

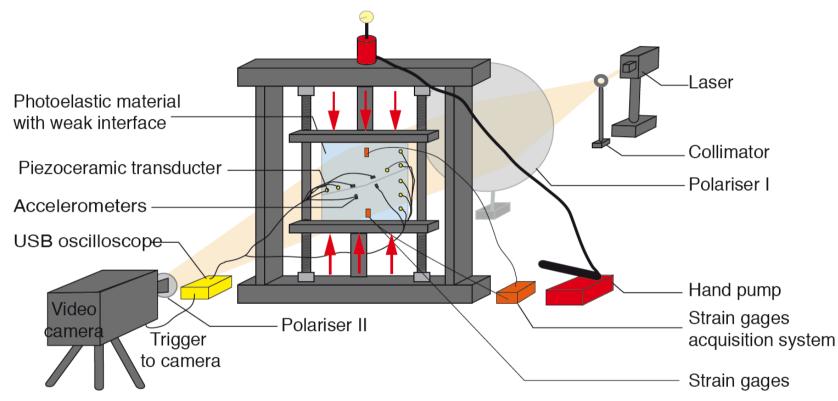
Quasi-dynamic ruptures

Nucleation and fracture propagation speed



- Rock fracture in the lab
- Precursory processes to Stick-slip
- Foreshock dynamics
- "Tidal" modulation

Stick-Slip in photo-elastic material: experimental set-up



Mode II dynamic rupture propagation on pre-cut, critically oriented, Columbia resin

(Nielsen et al. 2010)

Combining *high speed photo-elasticity* (rupture dynamics)

AND *high frequency accelerometry and acoustics* (near field strong motion and radiated wavefield) (Schubnel et al. 2011)

Propagation of a stick slip instability

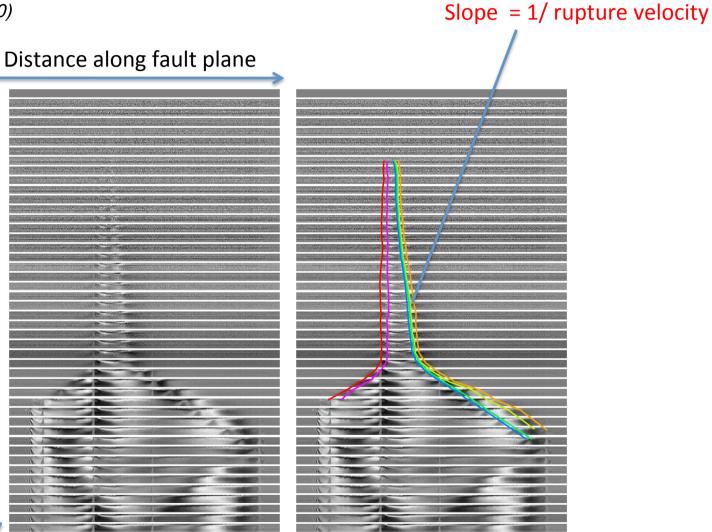
• Rupture propagation on fault using High speed camera



Rupture propagation speed of stick slip

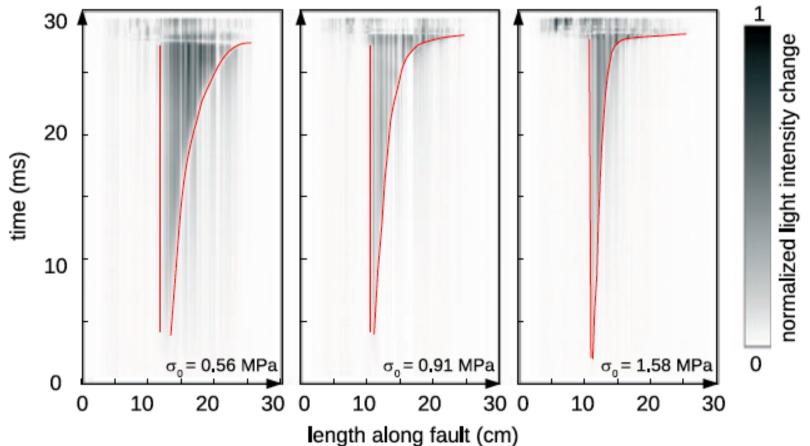
• Spontaneous ruptures studied using photoelasticity

(Nielsen et al. 2010)



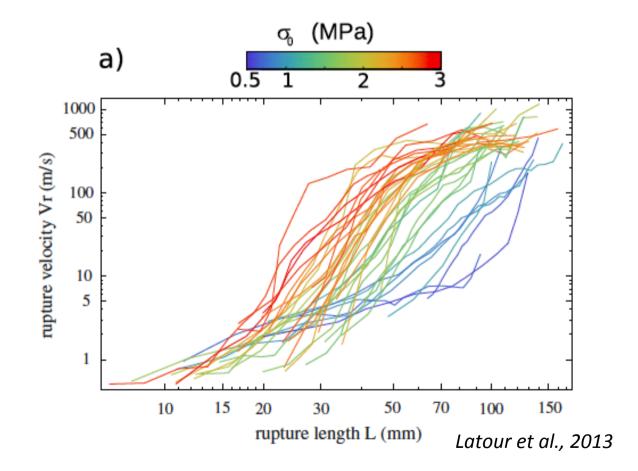
TIME

Nucleation of stick slip instabilities Influence of normal stress

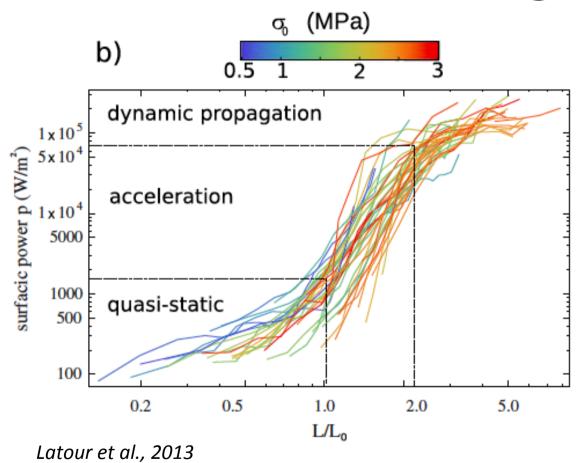


Latour et al., 2013

Nucleation of stick slip instabilities scaling

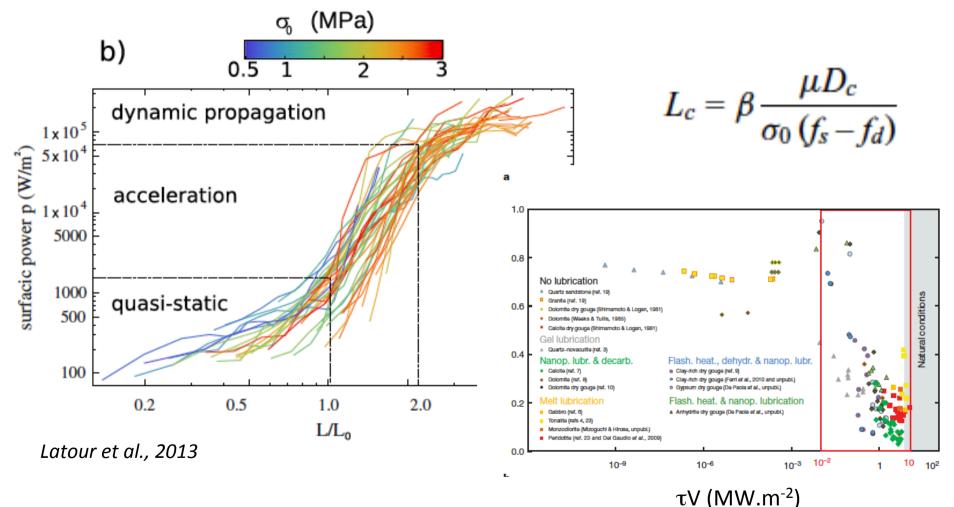


Nucleation of stick slip instabilities scaling



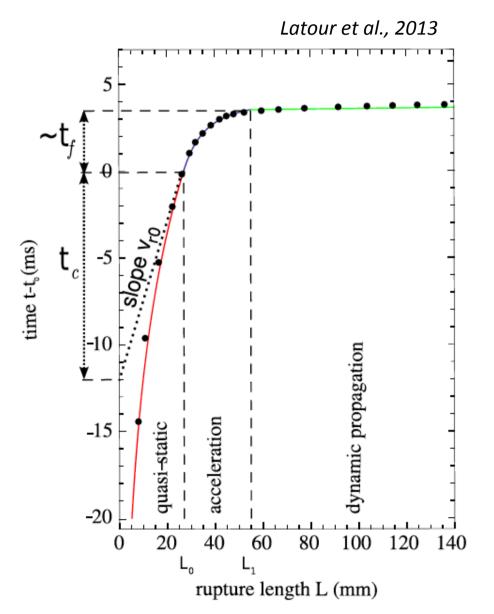
$$L_c = \beta \frac{\mu D_c}{\sigma_0 \left(f_s - f_d\right)}$$

Nucleation of stick slip instabilities scaling



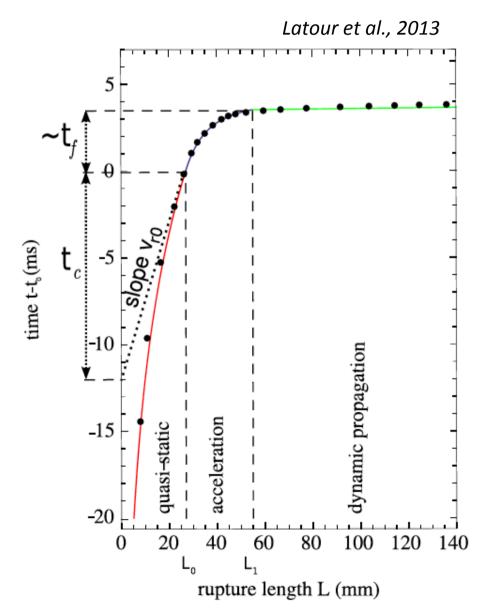
Di Toro et al., Nature 2011

Nucleation of stick slip instabilities: an exponential growth



$$t_c = L_0 / v_{r0}$$

Nucleation of stick slip instabilities: an exponential growth



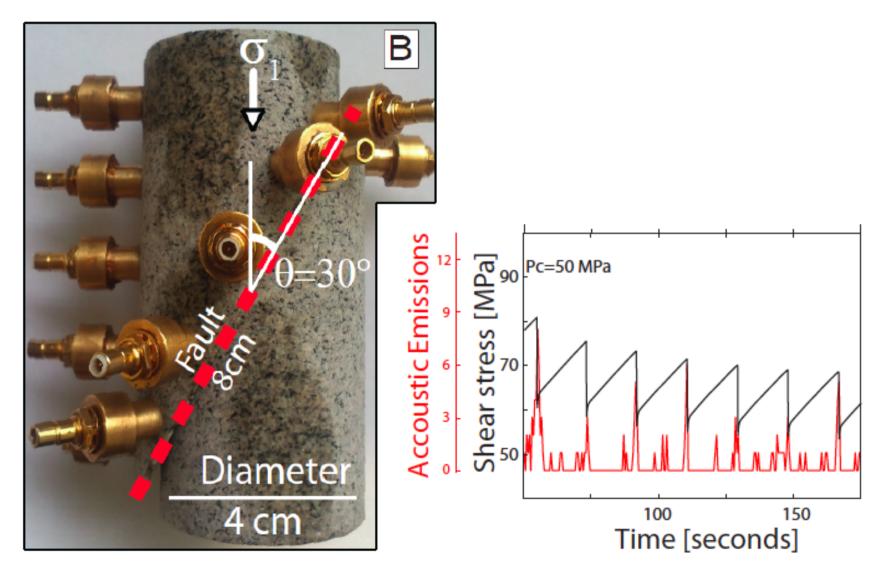
$$t_c = L_0 / v_{r0}$$

With k proportionality between rupture speed and sliding velocity ≈ 3x10³ (measured experimentally)

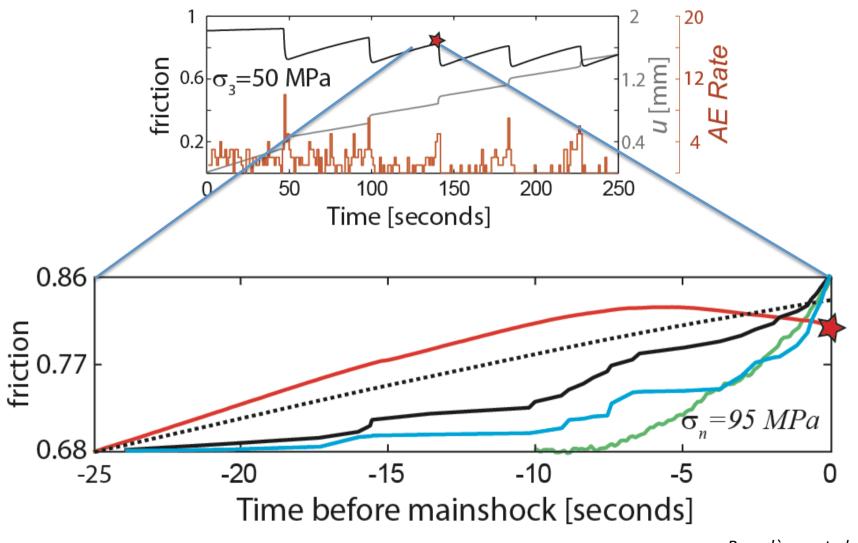
$$c_c = \frac{k\mu}{p_0} \frac{D_c}{f_s - f_d}$$

Stick-Slip in rocks (WG)

Saw cut Westerly granite – strain rates: 1-10 μ m/s

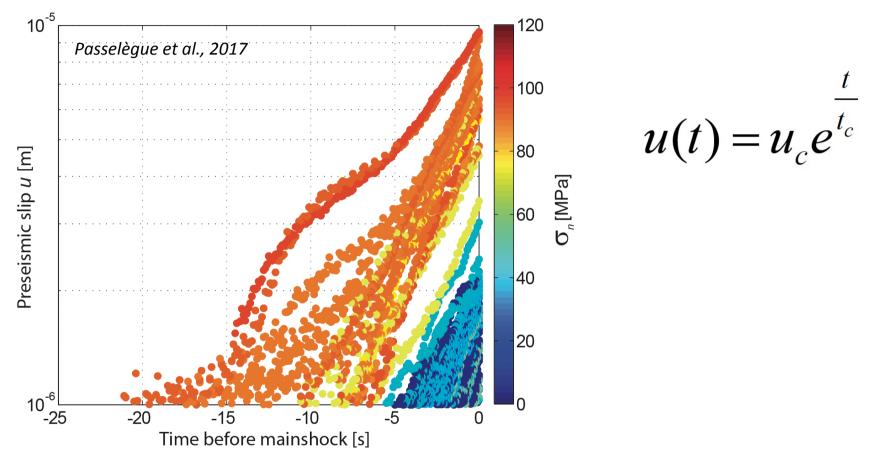


Precursories to Stick-Slip



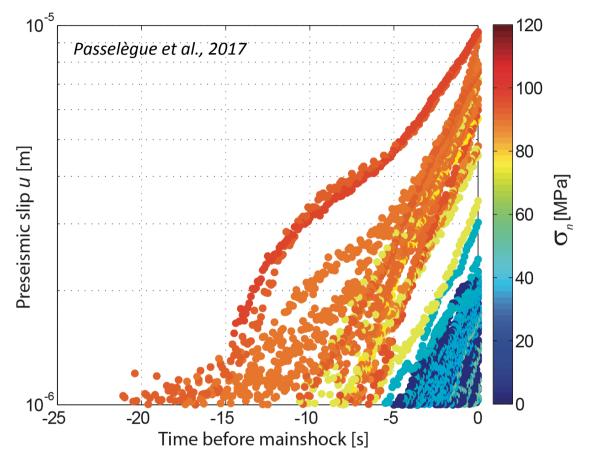
Passelègue et al., 2017

Precursory slip



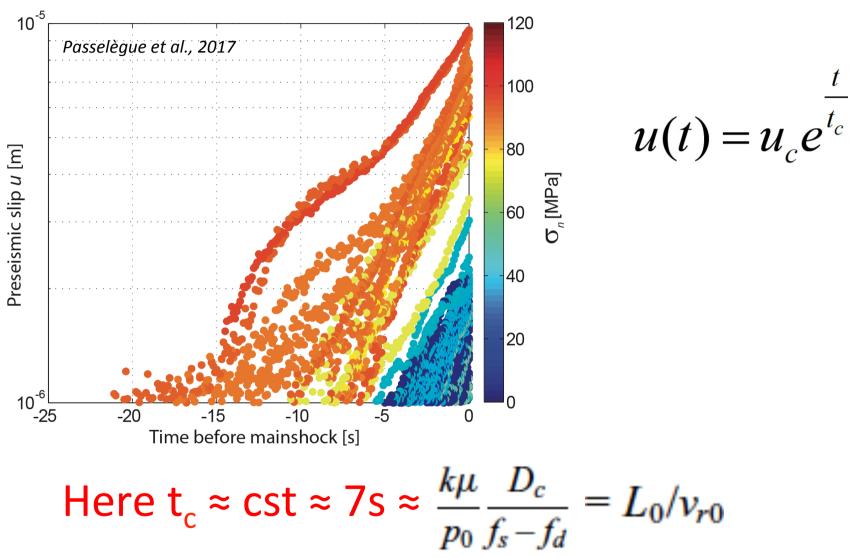
Precursory slip

 $u(t) = u_c e^{\overline{t_c}}$



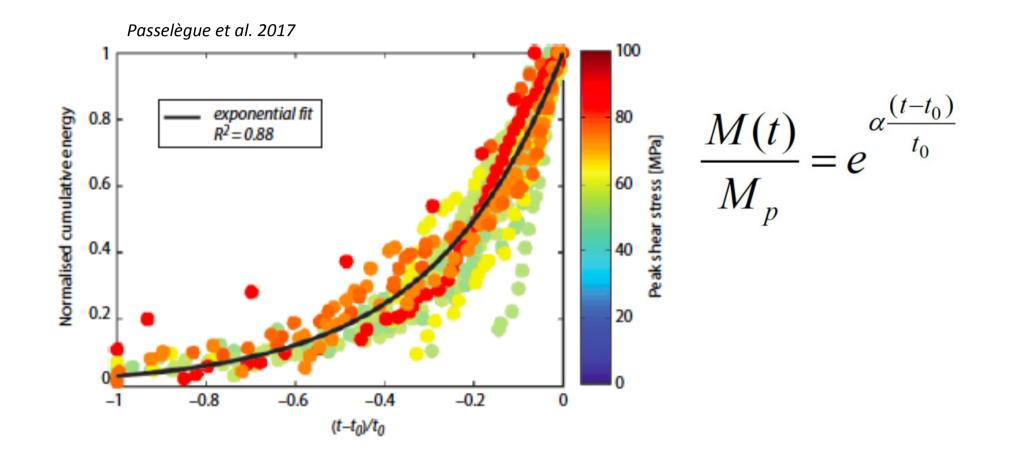
Here $t_c \approx cst \approx 7s$

Precursory slip

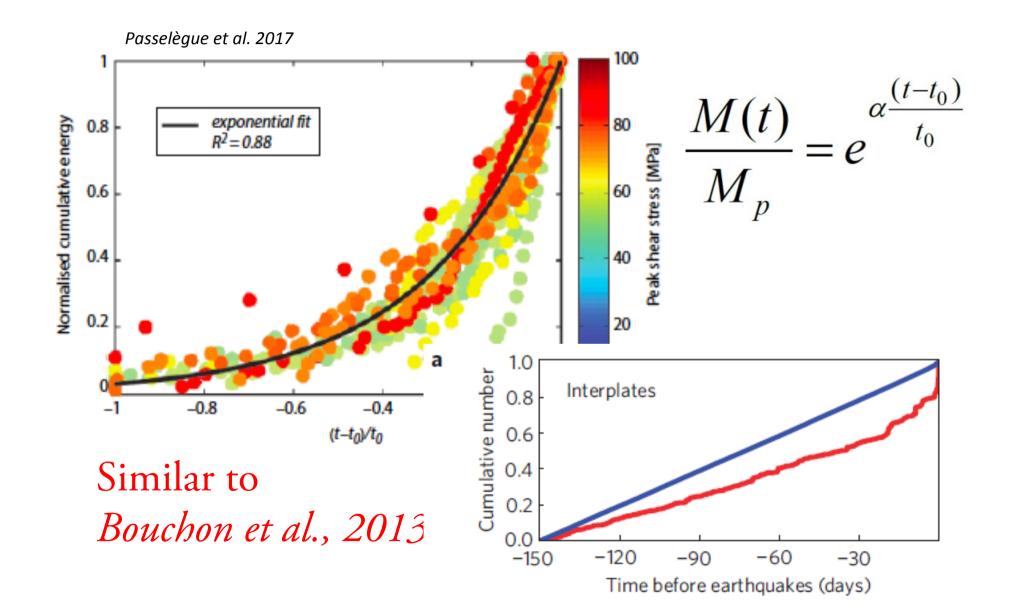


Latour et al., 2013

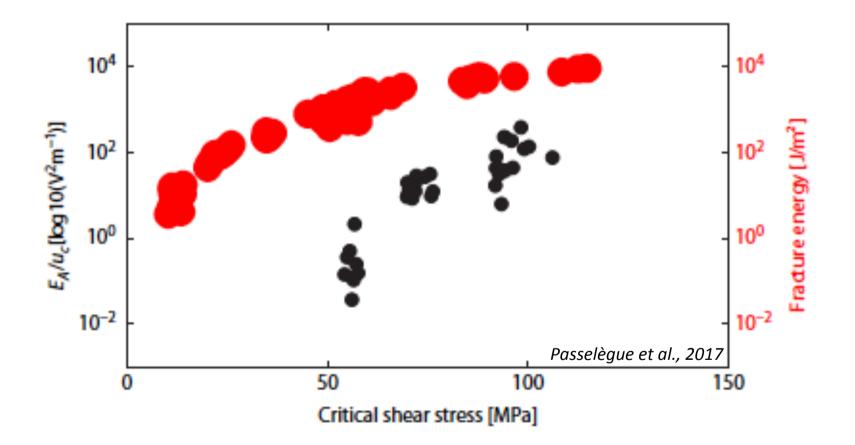
Foreshock dynamics



Foreshock dynamics

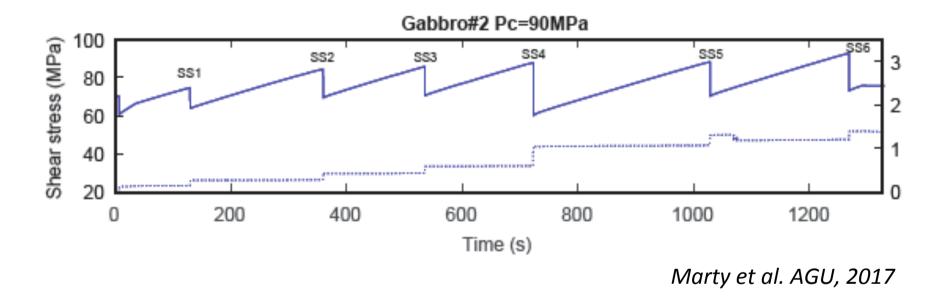


Evolution of seismic/aseismic coupling with fracture energy G & shear stress

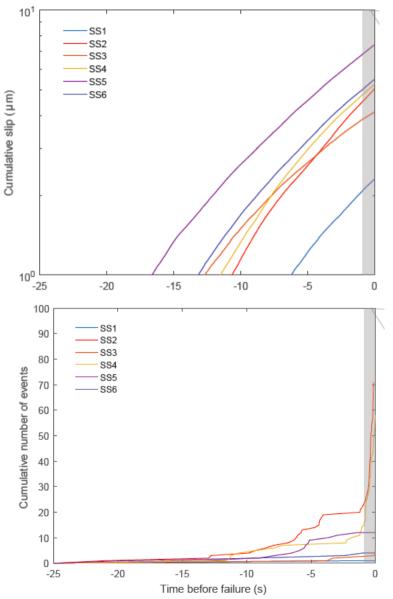


- Rock fracture in the lab
- Precursory processes to Stick-slip
- Foreshock dynamics
- "Tidal" modulation

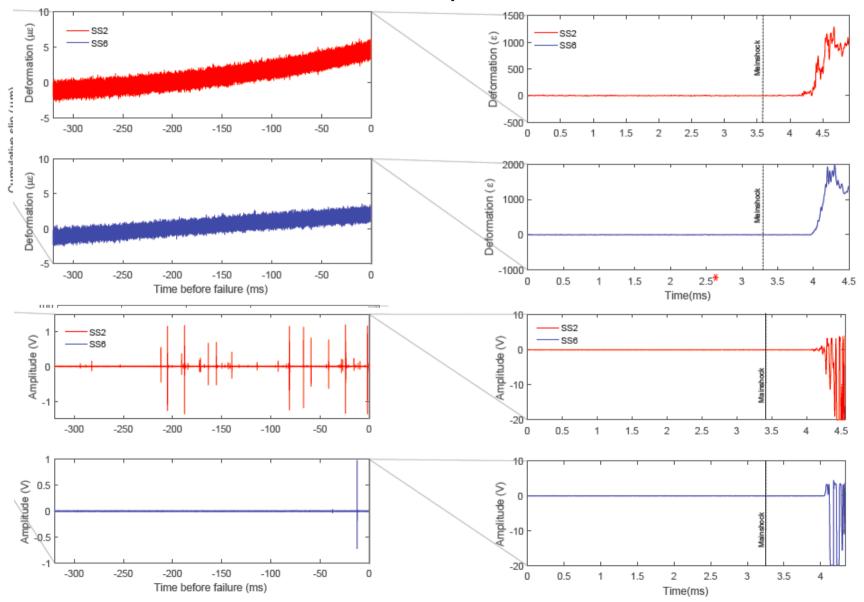
• Continuous record of stick-slip failure in Gabbro



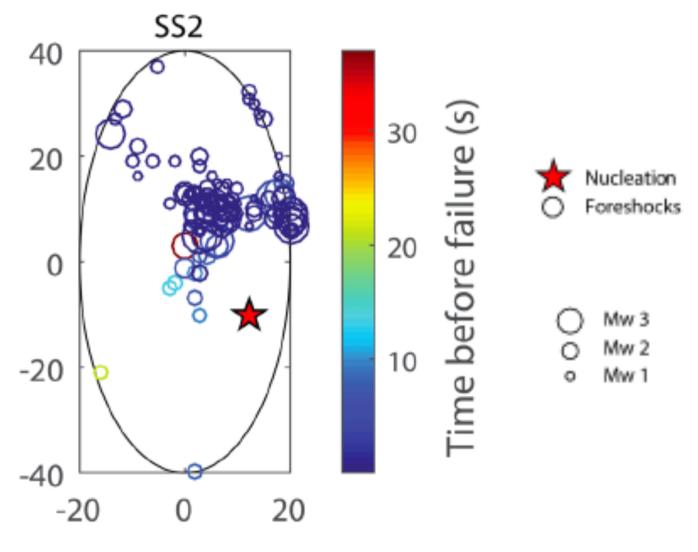
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• Continuous record of stick-slip failure in Gabbro

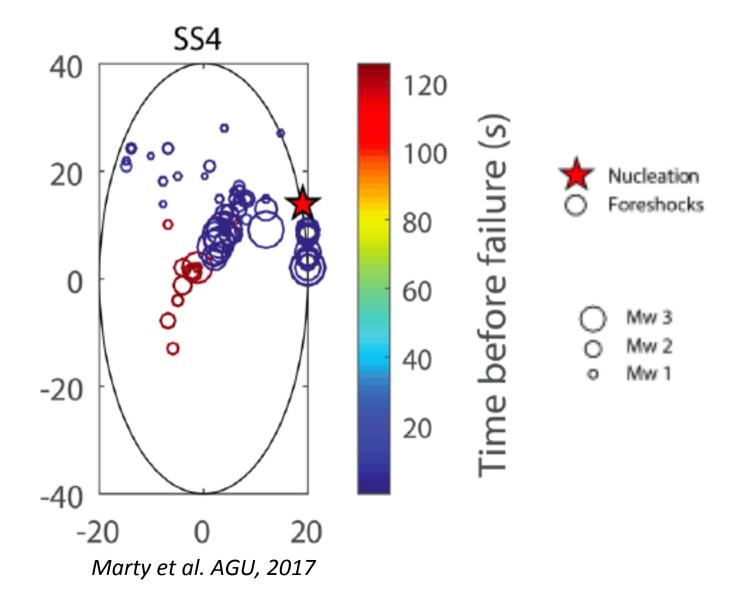


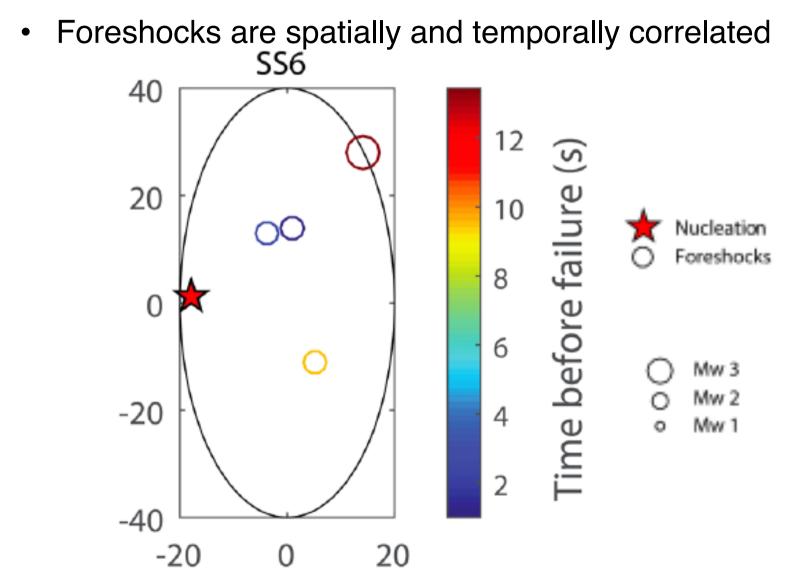
• Foreshocks are spatially and temporally correlated



Marty et al. AGU, 2017

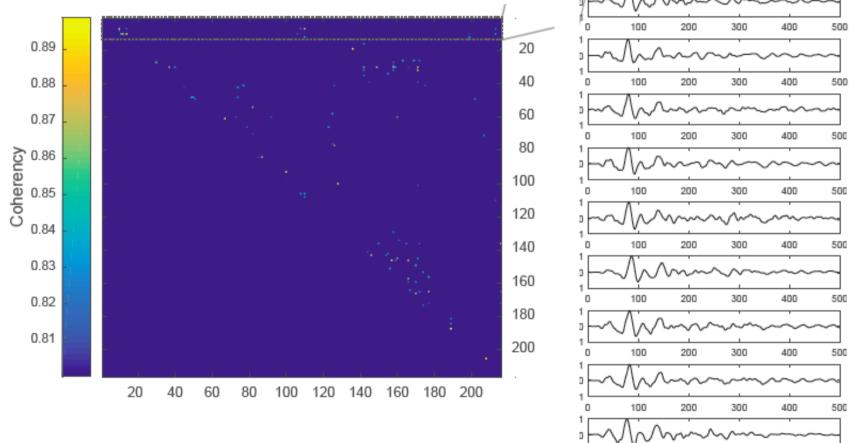
• Foreshocks are spatially and temporally correlated





Marty et al. AGU, 2017

Life of a single asperities during and over several cycles
All foreshocks



0

100

200

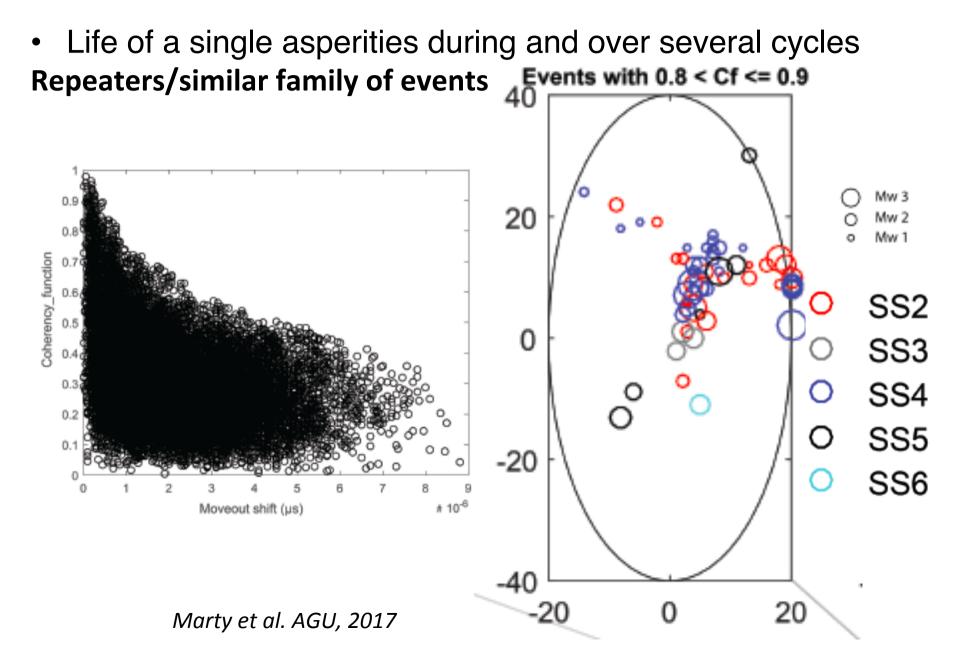
Time (µs)

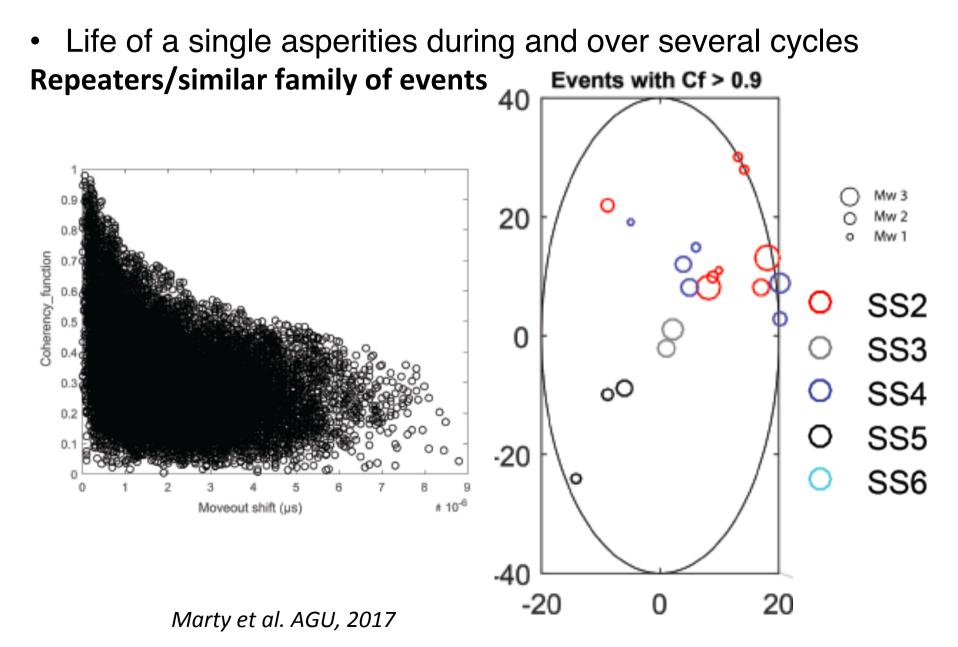
300

400

500

Marty et al. AGU, 2017



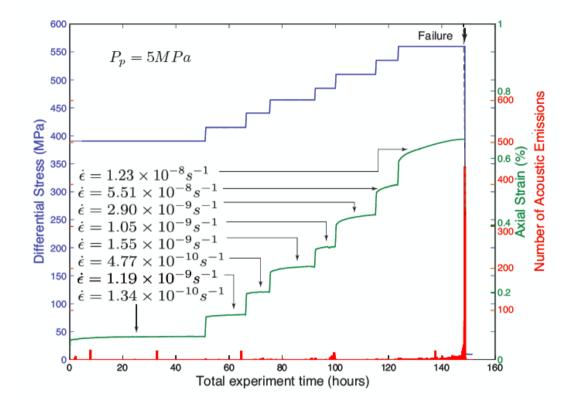


MOTIVATION

- Rock fracture in the lab.
- Precursory processes to Stick-slip
- Foreshock dynamics
- "Tidal" modulation

How to simulate tidal modulations?

Typical Stress and Strain functions



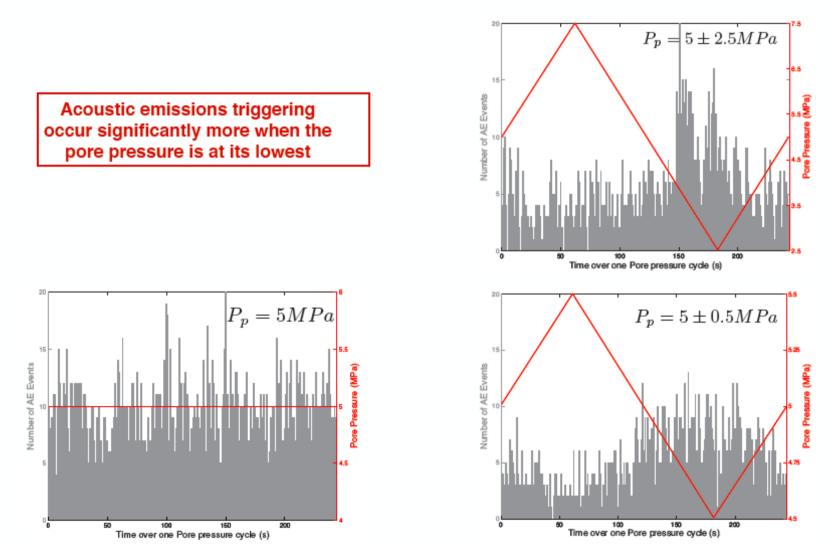
- Set of triaxial deformation experiments on water-saturated Fontainebleau sandstones, 7% porous, T=35°C, Pc=35MPa
- Samples loaded by:
 - constant stress (creep)

 tectonic loading
 - sinusoidal pore pressure variations --> tidal or seasonal

loading

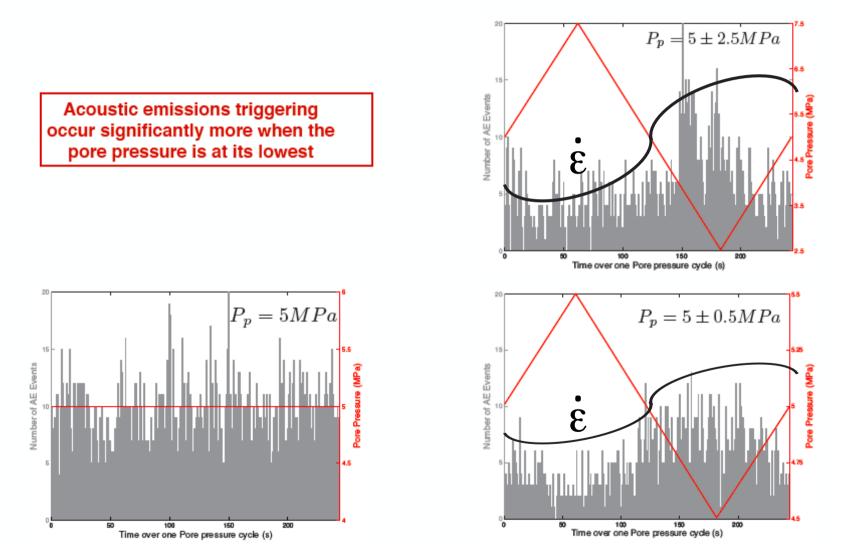
GLOBAL AVERAGE (over one Pp cycle) of AE catalog





GLOBAL AVERAGE (over one Pp cycle) of AE catalog

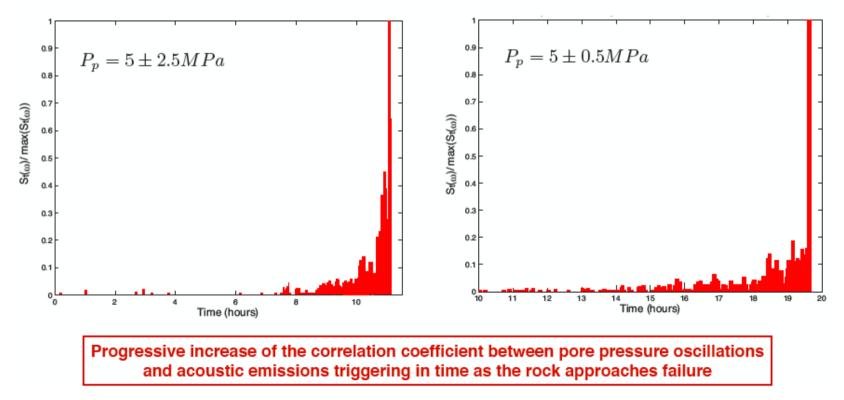
Experimental Results: Acoustic Emissions during failure stress step



A growing stress suceptibility (i.e. sensitivity)

Experimental Results: Acoustic Emissions triggering during failure stress step

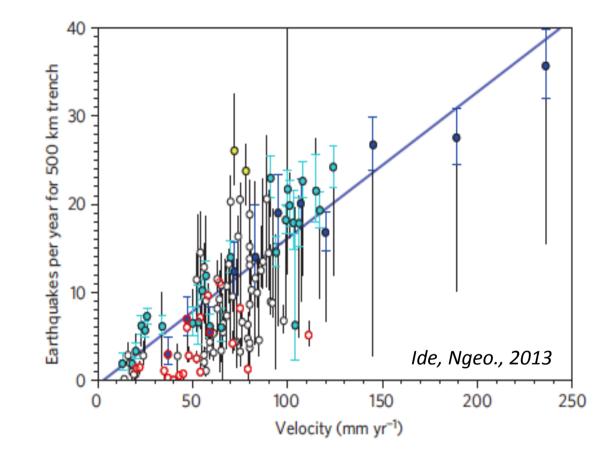
Correlation coefficient between AE and pore pressure:



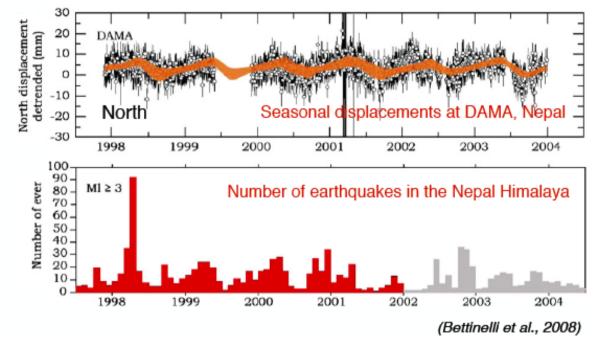
 $S_t(\omega) = f($ pulsation of the pore pressure oscillations, AE rate)

Chanard 2015; Chanard et al. in prep.

The proportionality between relative plate velocity and seismicity in subduction zones



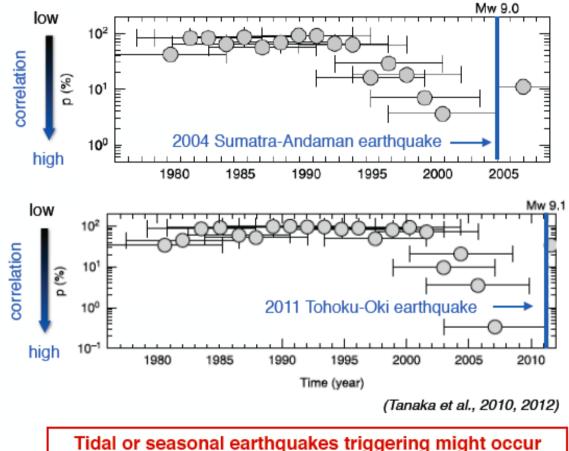
Seasonal seismicity in the Himalayas



- Seismicity rate in the winter twice as high as in the summer in the Nepal Himalaya
- Other regional observations of seismicity rate with periodic loading (tidal, seasonal)
- No global systematic correlation of seismicity with amplitude or frequency of periodic loading

Tidal triggering during large EQ nucleation

Observation: increase of correlation between seismicity and oceanic tides prior to large earthquakes



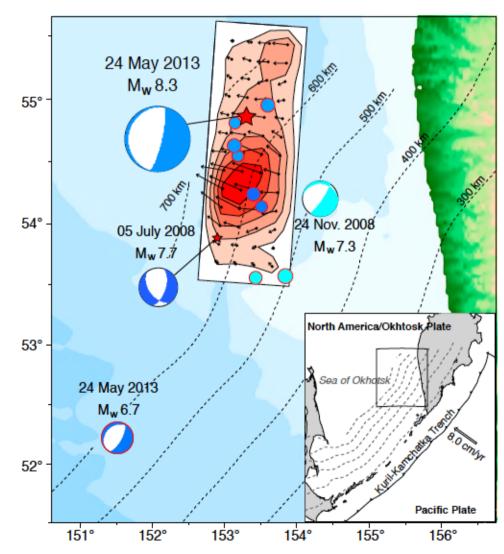
favorably during the long nucleation phase

Conclusions

- During rock fracture, Omori's law is observed for foreshocks and aftershocks, and small stress perturbations can modulate seismicity rate. Modulation is correlated with strain rate and thus suceptibility grows as failure approaches.
- Nucleation phase of stick slip instability corresponds an exponential growth of rupture length & slip, and thus slip velocity & rupture speed. Foreshock and aseismic moment releases are also exponential. Seismic/aseismic coupling increases with fracture energy and normal stress.
- Nucleation phase is also charaterized by repeaters (repetitive failure of the same asperity).

Do large deep earthquake...

The Okhotsk, May 24th 2013, Mw=8.3, 620km deep EQ

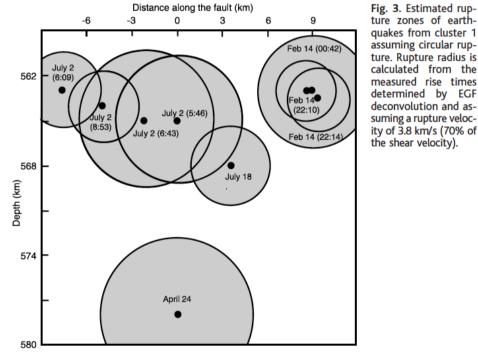


Ye et al. Science 2013

...have a nucleation phase? **Repeating Deep Earthquakes: Evidence for Fault Reactivation**

at Great Depth

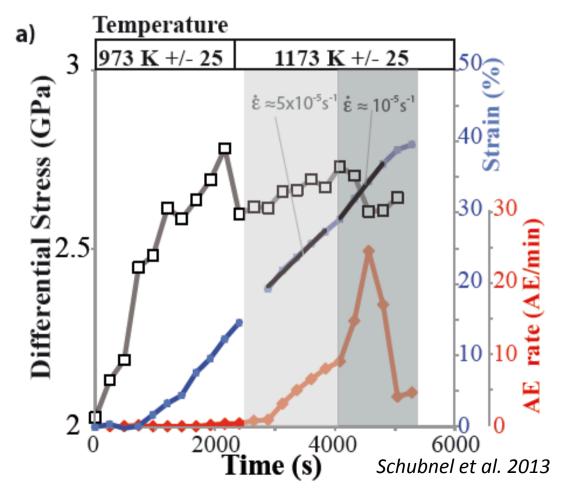
Douglas A. Wiens and Nathaniel O. Snider



ture zones of earthquakes from cluster 1 assuming circular rupture. Rupture radius is calculated from the measured rise times determined by EGF deconvolution and assuming a rupture velocity of 3.8 km/s (70% of the shear velocity).

Nature 2001

Sintered Mg₂GeO₄ – 30µm grain size Effective mean stress $(\sigma_1 + 2\sigma_3)/3 = 4$ GPa +/-0,25 Strain rate = 10⁻⁴/s

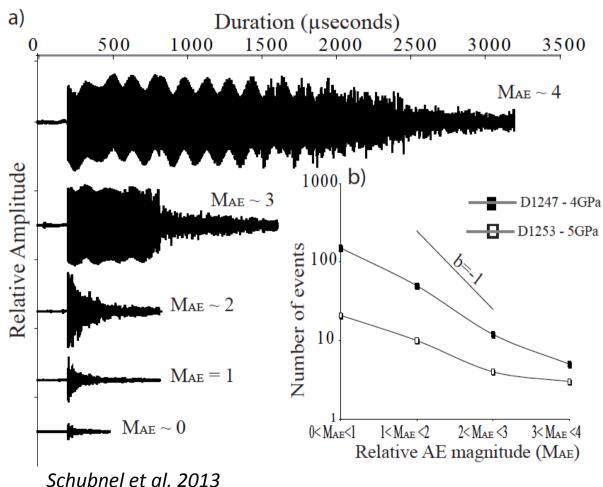


Two complete AE catalogues

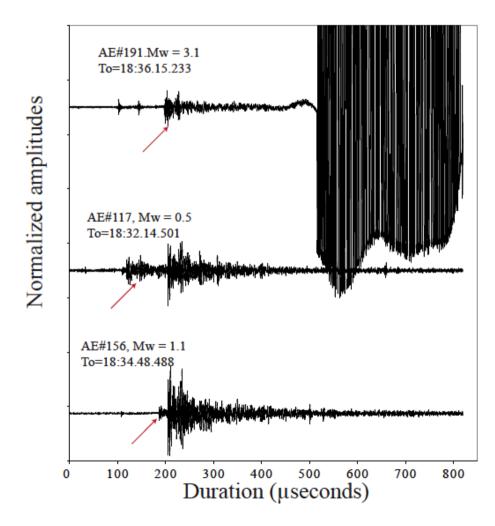
D1247 Effective mean stress $(\sigma_1+2\sigma_3)/3 = 4$ GPa +/-0,25 D1253 Effective mean stress $(\sigma_1+2\sigma_3)/3 = 5$ GPa +/-0,25 Strain rate = 10⁻⁴/s



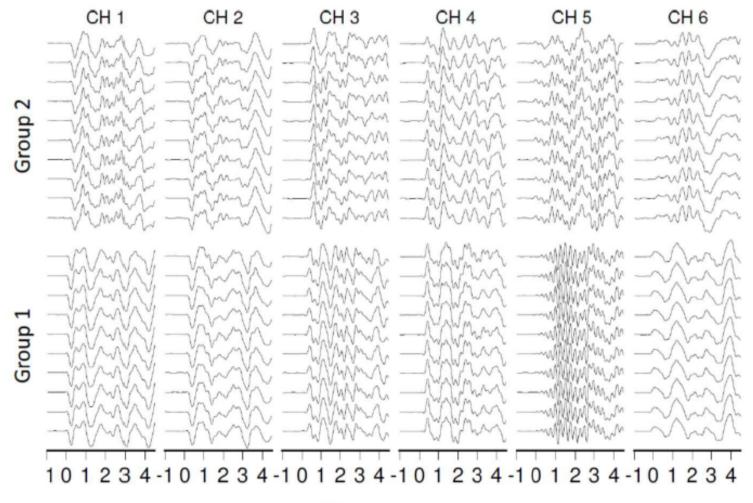
Sonification: courtesy to Ben Holtzman LDEO, U. Columbia NY Visiting prof. ENS 2015



Looking closer into the AE source - *Double source or nucleation phase?* Approximately 20% of the AEs have a complex source functions



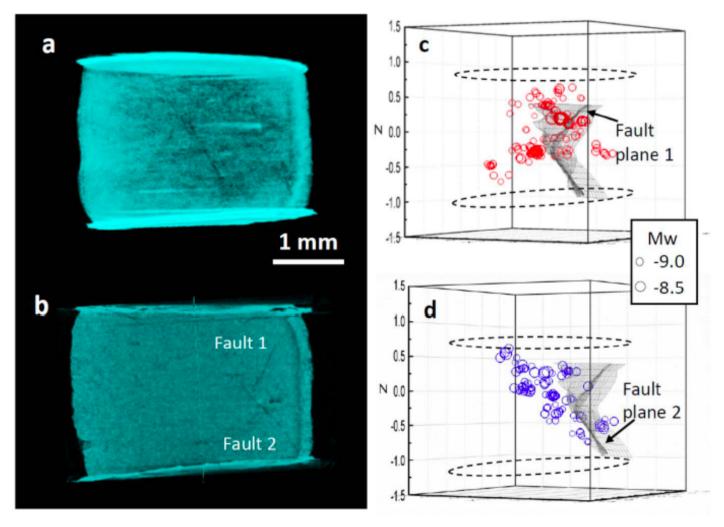
Double difference relocation (Waldhauser and Ellsworth 2000)



Time, µs

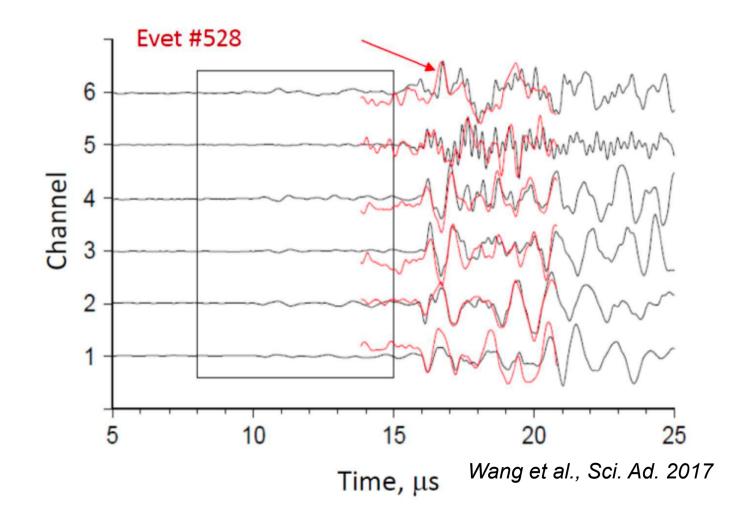
Wang et al., Sci. Ad. 2017

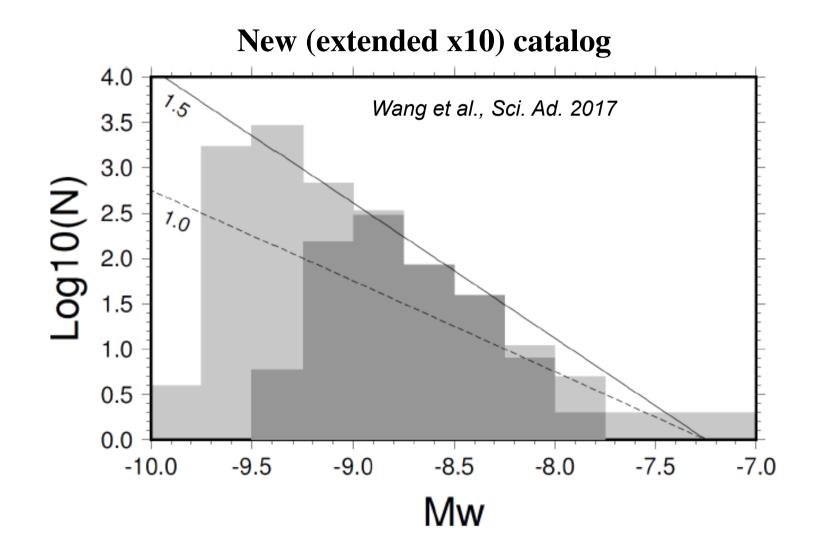
Correlating X-ray tomography and AE relocations



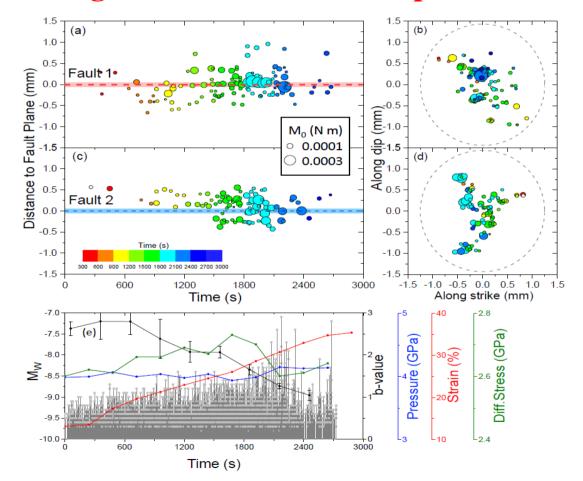
Wang et al., Sci. Ad. 2017

Template matching of continuous wfms





Signature of a nucleation phase?



Wang et al., Sci. Ad. 2017